

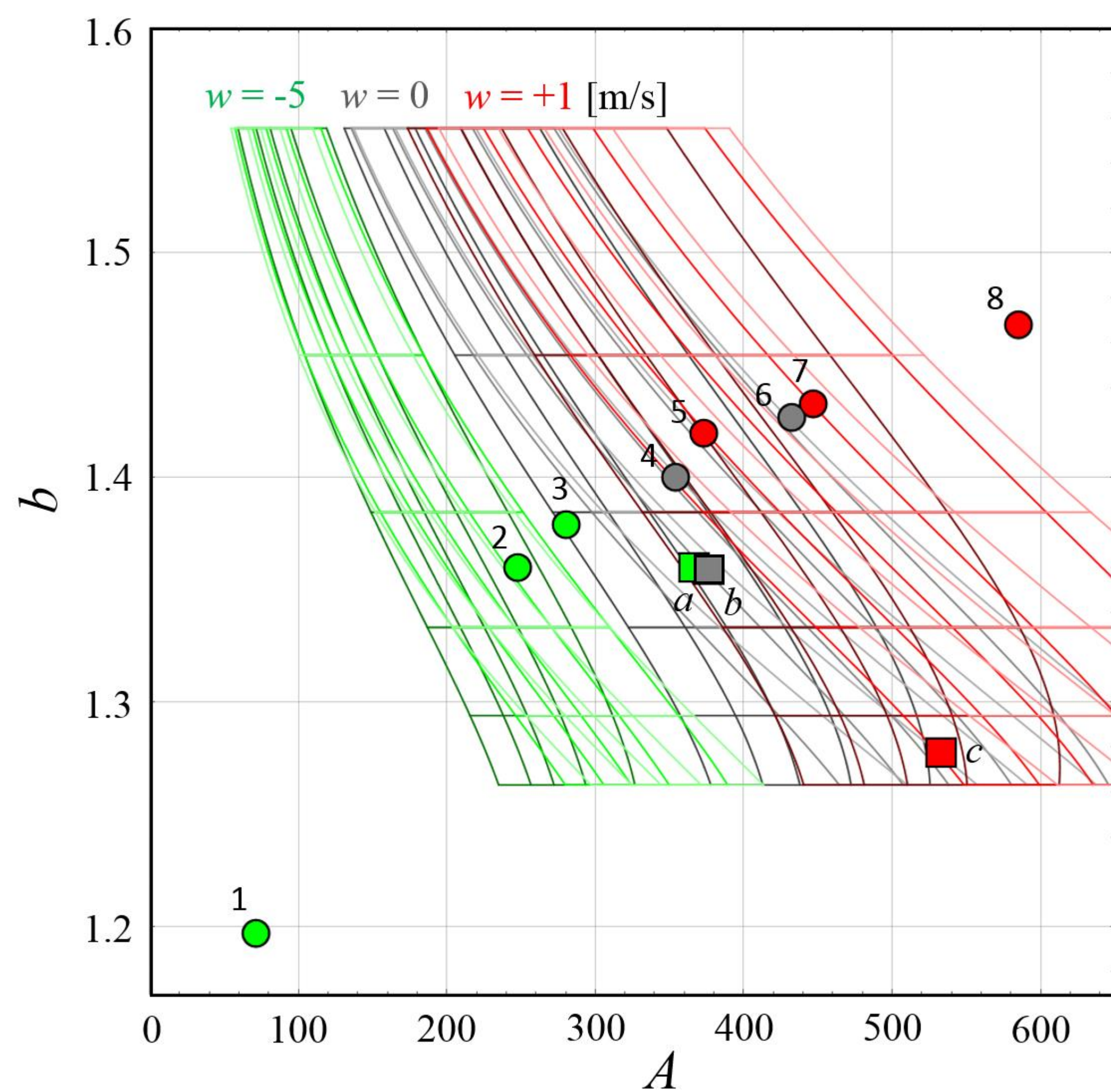
# A Possible Explanation for the Z-R Parameter Inconsistency when Comparing Stratiform and Convective Rainfall



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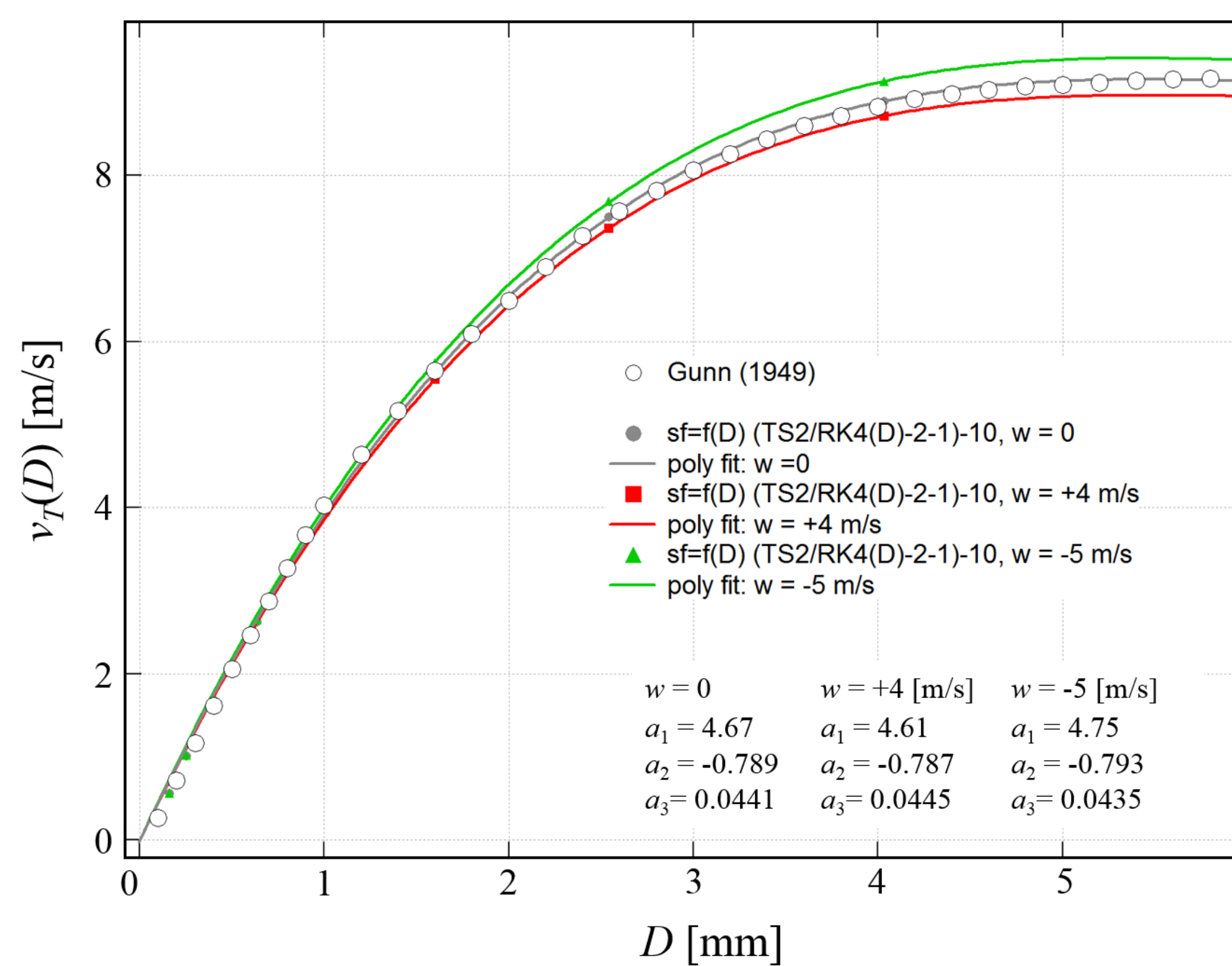
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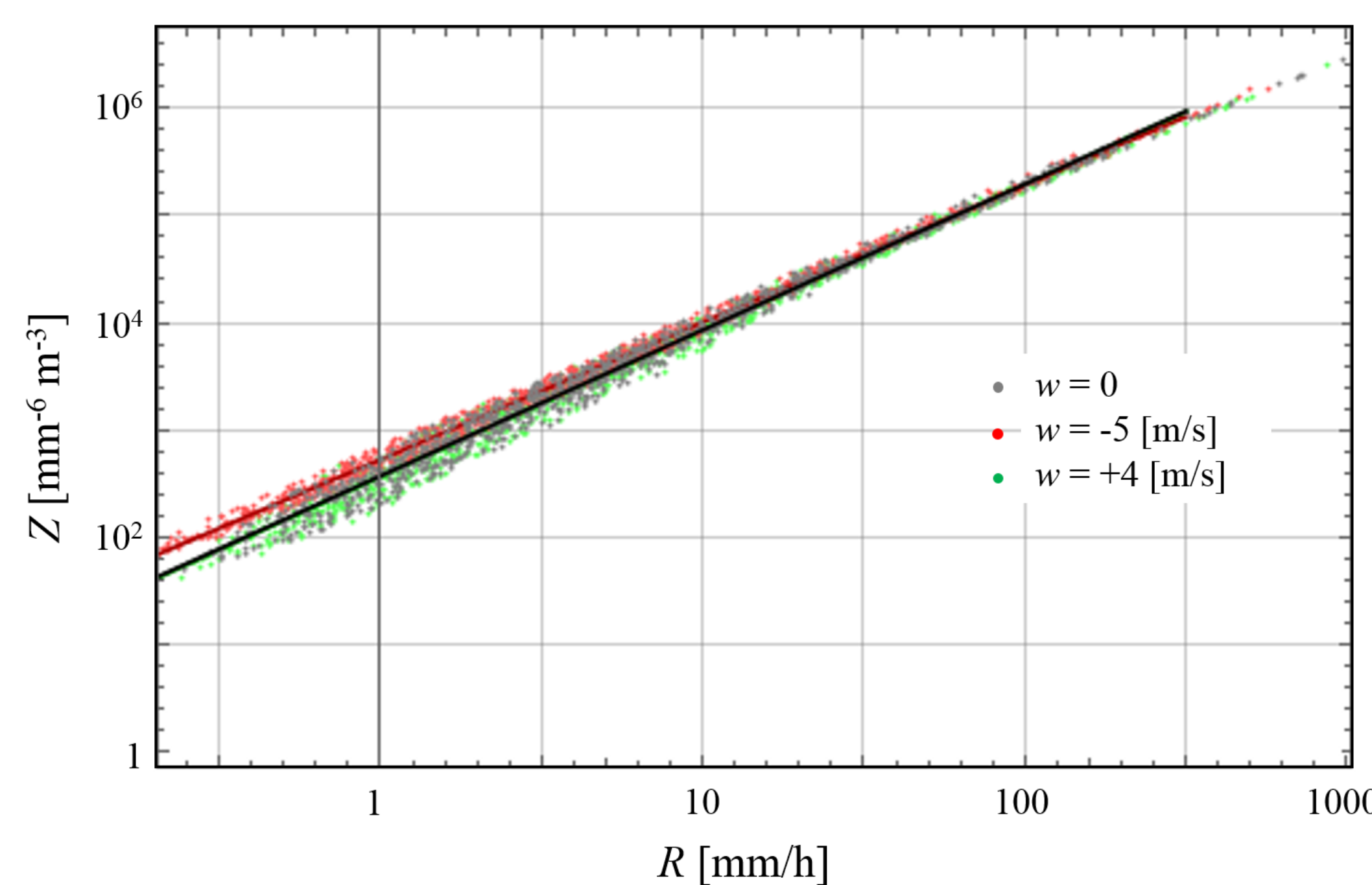
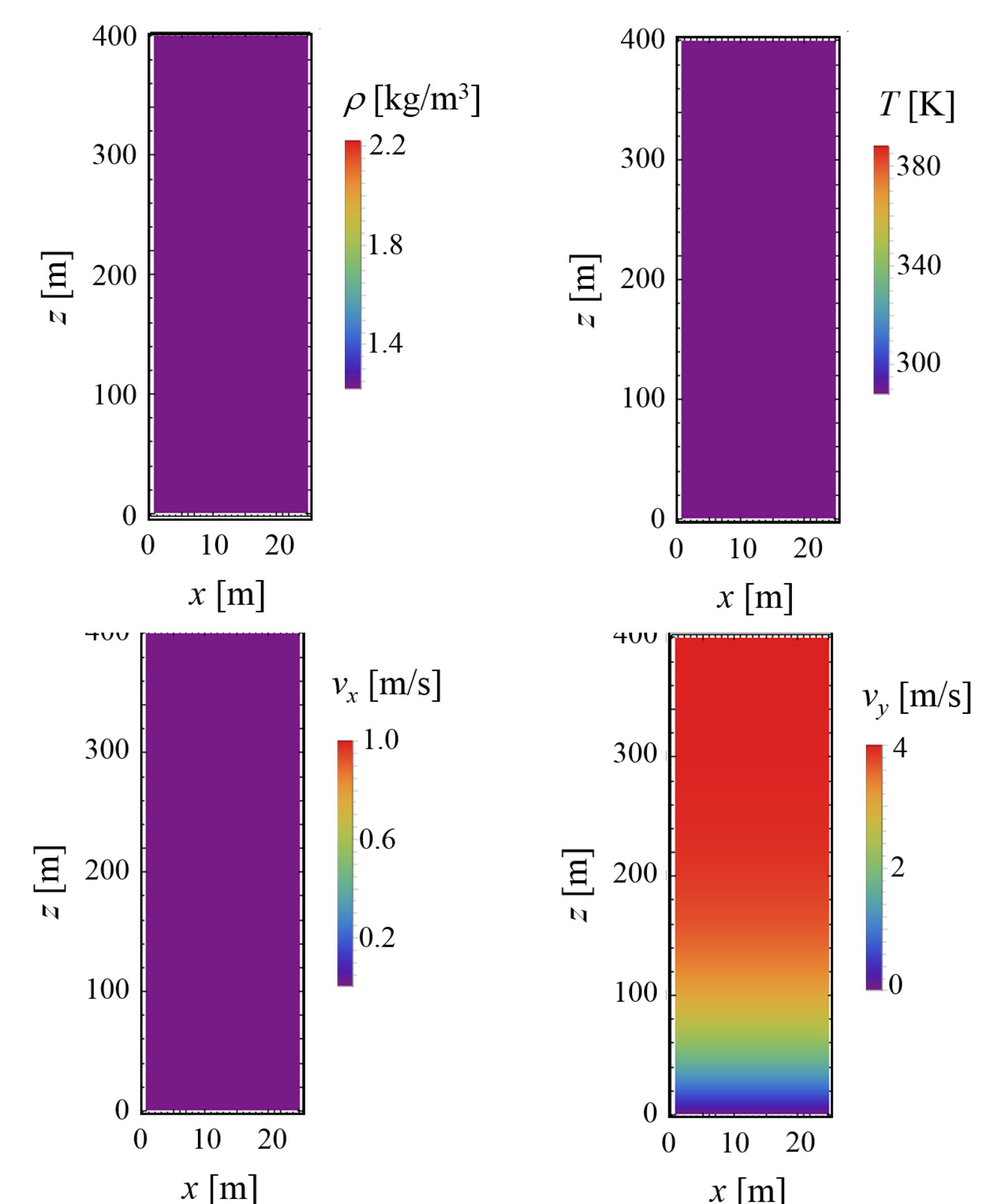
Circles are JWD derived  $A$ - $b$  pairs (near Athalassa, Cyprus) from selected 24 hour averages during 2011 - 2014. Gray circles represent stratiform events defined by rainfall rates that did not exceed 10 mm/h at anytime during the 24-hour period. Green and red circles represent convective events defined by rainfall rates greater than or equal to 10 mm/h at anytime during the 24 hour period. Green circles are  $A$ - $b$  pairs that fall to the left of the stratiform gray circles, while and red circles are pairs that fall to the right. The lines are output from a model simulation using the 3-parameter gamma DSD with drop terminal velocity approximated as  $v(D) = \alpha D^\beta - w$ . For various values of  $N_0$ ,  $\mu$ , and  $\Lambda$ , gray lines correspond to  $w = 0$ , green lines to  $w = -4$  [m/s], and red lines to  $w = +1$  [m/s]. The squares correspond to the  $A$ - $b$  pairs from the particle trajectory simulation in the graph to the right.

The well-known Z-R power law  $Z = AR^b$  uses two parameters,  $A$  and  $b$ , in order to relate rainfall rate  $R$  to measured weather radar reflectivity  $Z$ . A common method used by researchers is to compute  $Z$  and  $R$  from disdrometer data and then extract the  $A$ - $b$  parameter pair from a log-linear line fit to a scatter plot of Z-R pairs. Even though it may seem far more truthful to extract the parameter pair from a fit of radar  $Z_R$  versus gauge rainfall rate  $R_G$ , the extreme difference in spatial and temporal sampling volumes between radar and rain gauge creates a slew of problems that can generally only be solved by using rain gauge arrays and long sampling averages. Disdrometer derived  $A$ - $b$  parameters are easily obtained and can provide information for the study of stratiform versus convective rainfall. However, an inconsistency appears when comparing averaged  $A$ - $b$  pairs from various researchers. Values of  $b$  range from 1.26 to 1.51 for both stratiform and convective events. Paradoxically the values of  $A$  fall into three groups: 150 to 200 for convective; 200 to 400 for stratiform; and 400 to 500 again for convective. This apparent inconsistency can be explained by computing the  $A$ - $b$  pair using the gamma DSD coupled with a modified drop terminal velocity model,  $v(D) = \alpha D^\beta - w$ , where  $w$  is a somewhat artificial constant vertical velocity of the air above the disdrometer. This model predicts three regions of  $A$ , corresponding to  $w < 0$ ,  $w = 0$ , and  $w > 0$ , which approximately matches observed data.



Using CFD based single particle trajectory modeling and a vertical wind profile, the terminal velocity can be approximated under any given wind conditions. The resultant terminal velocity function is then approximated by a 3<sup>rd</sup> order polynomial. Three cases are simulated:  $w = 0$ ,  $w = -5$  [m/s], and  $w = +4$  [m/s].  $A$ - $b$  parameter pairs are then found by performing a Monte Carlo Z-R scatter plot using the three-parameter gamma DSD.

**CFD Input to Particle Trajectory Code:** air density, air temperature, horizontal air velocity, and vertical air velocity. Below are the inputs for the  $w = +4$  case.



Joss Disdrometer Data  
Anthalassa, Cyprus,  
35.15N, 33.40W



1: Aug 10, 2011

2: May 29, 2011

3: Apr 17, 2013

4: Jan 27, 2012

5: Oct 18, 2013

6: Sep 23, 2011

7: Oct 24, 2012

8: May 10, 2013

